

vae of the phantom midge *Chaoborus* (Chaoboridae) and cyclopoid copepods (Kerfoot, 1975; Arnott and Vanni, 1993; Schilling *et al.*, 2009). *Chaoborus* are gape-limited ambush predators (Spitze, 1985) that are efficient at consuming small zooplankton prey (Mumm, 1997). Chaoborid community composition is a significant determinant of zooplankton community structure, although predation has been found to play less of a role than environmental factors in structuring zooplankton communities in relatively small and shallow Boreal Shield lakes (Kurek *et al.*, 2011). Copepods also exert a strong predation effect upon zooplankton, however their historical abundance is difficult to ascertain, as they leave poorly preserved remains in the sediment record (Frey, 1964).

Historic predation regimes can be reconstructed by examining the fossil remains of organisms within lake sediments (Jeppesen *et al.*, 2001; Palm *et al.*, 2005; Sweetman and Smol, 2006). Fish leave few remains in the sediment record, thus their population dynamics, introduction, or extirpation must often be inferred indirectly from proxy evidence. A major zooplankton predator, *C. americanus* is well-suited as an indirect indicator of fishless conditions (Sweetman and Smol, 2006) due to its lack of DVM in the water column, large size, and pigmentation, which make it susceptible to planktivorous fish predation (Pope *et al.*, 1973; von Ende, 1979; Wissel *et al.*, 2003). Additionally, third and fourth larval instar chaoborids leave well-preserved and readily identifiable remains in the sediment record (Uutala, 1990). As a result, *C. americanus* has been widely utilised in North America to indirectly track fish introductions and extirpations through time (Lamontagne and Schindler, 1994; Uutala *et al.*, 1994; Schilling *et al.*, 2008). Non-DVM chaoborids in western Europe, such as *C. obscuripes*, are also used to determine historic fish population dynamics from sediment records (Palm *et al.*, 2005; Luoto and Nevalainen, 2009; Palm *et al.*, 2012). The cladoceran *Bosmina*, a key herbivore and cosmopolitan zooplankton taxon, also responds to the dominant predation regime within lakes and have been utilised as an indirect indicator of predation (Palm *et al.*, 2005; Alexander and Hotchkiss, 2010; Korosi *et al.*, 2010, 2013). Several *Bosmina* size attributes, including antennule and mucro length, can vary significantly based upon the dominant predation regime (Kerfoot, 1975). Additionally, *Bosmina* leaves numerous, well-preserved remains in the sediment record, allowing statistically-valid size measurements to be obtained from sediment intervals (Brahney *et al.*, 2010). Therefore, the examination of both *Chaoborus* and *Bosmina* remains from lake sediment records provides an opportunity to assess predation regime shifts across longer periods of time than monitoring data sets or mesocosm studies (when available) allow.

Here, we examine the impacts of a 1980s minnow

species introduction and also centennial-scale shifts in predation intensity, within a historically fishless lake near Sudbury, Ontario, Canada. The minnow introduction provided a *natural experiment* to assess the effects of a relatively recent fish introduction compared to historic shifts in predation intensity. By incorporating the temporal dynamics of a major invertebrate predator (*i.e.* *Chaoborus*), as well as its common prey (*i.e.* *Bosmina*), we build upon other paleolimnological studies of historic shifts in zooplankton size structure from eastern North America (Leavitt *et al.*, 1989; Korosi *et al.*, 2010, 2012). Whereas most paleolimnological studies recognize the importance of top-down controls of predation on zooplankton (*i.e.* fish introduction, extirpation, and population shifts), climate-induced shifts in lake thermal structure (Magnuson *et al.*, 1997; Keller, 2007; Quinlan *et al.*, 2012) also have the potential to significantly alter chaoborid-zooplankton dynamics due to temperature-dependent processes (Macphee *et al.*, 2011). This is important to consider, given that, within our study region, 20th century temperature and precipitation patterns are largely different from those of previous centuries (Buhay and Edwards, 1995; Magnuson *et al.*, 1997; Buckley *et al.*, 2004).

METHODS

Site description and field methods

The study lake (W16) was selected from among lakes monitored as part of the Environment Canada's Acid Rain Biomonitoring Program (McNicol *et al.*, 1995, 1996). In early-October 2010, a 35-cm long sediment core was obtained from W16 (46°53'3.83" N, 80°49'32.15" W). W16 is a remote, small (4.4 ha), shallow (~8 m maximum depth) headwater lake ~45 km northeast of Sudbury, Ontario (Fig. 1). In mid-August of 2005, the lake was thermally stratified (temperatures ranged from 23.4°C at 0.5 m to 8.7°C at 7.5 m), and anoxic [<1 mg L⁻¹ dissolved oxygen (DO)] below 3 m water depth. October 2010 water chemistry measurements indicate that W16 is slightly acidic (pH=6.6), oligotrophic (total phosphorus=8.4 µg L⁻¹), and has a relatively stained water colour (DOC=6.50 mg L⁻¹) (Fig. 2). The core was collected from the center of the lake using a gravity corer (Glew, 1989). Sediments were sectioned in the field at intervals of 0.5-cm resolution below 5 cm, and 0.25-cm resolution from 5 cm to the surface sediments using a vertical extruder (Glew, 1988). Near-annual water chemistry measurements have been obtained at W16 since 1983 (Fig. 2), and fish populations were also monitored once every few years since 1983 (Tab. 1). Minnow trap data indicate that small numbers of minnows may have been present in W16 during the 1980s, but populations dominated by northern redbelly dace (*Phoxinus eos*) became well established by the mid-1990s (Tab. 1).

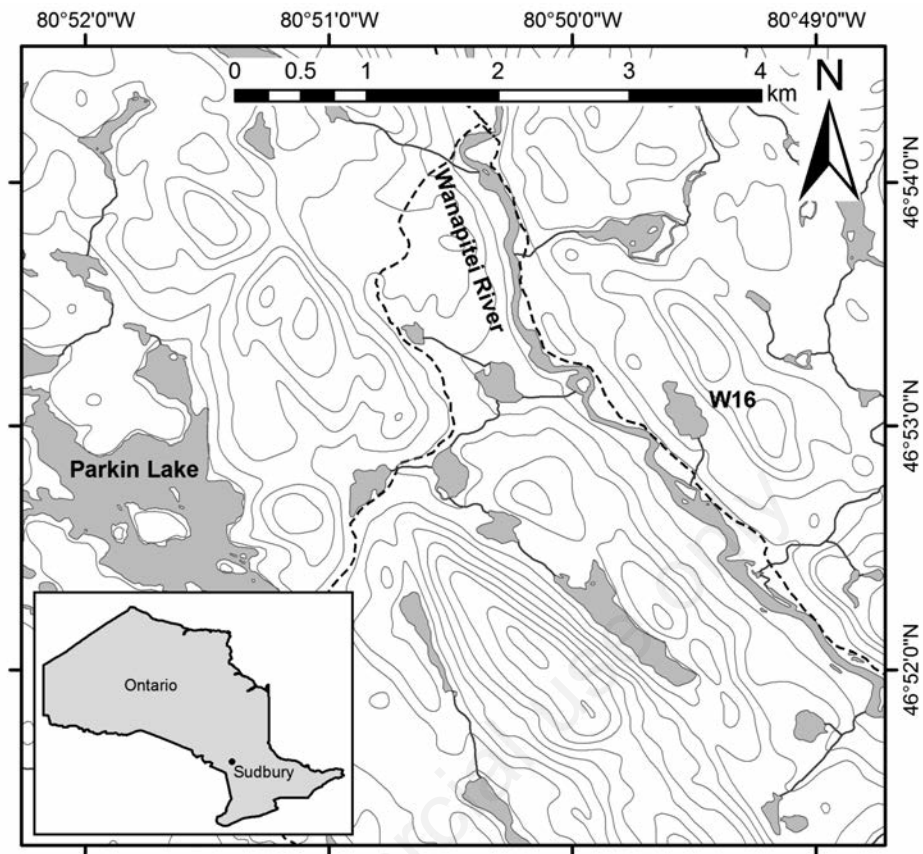


Fig. 1. Maps showing the locations of W16 and the location of the City of Greater Sudbury within the province of Ontario, Canada.

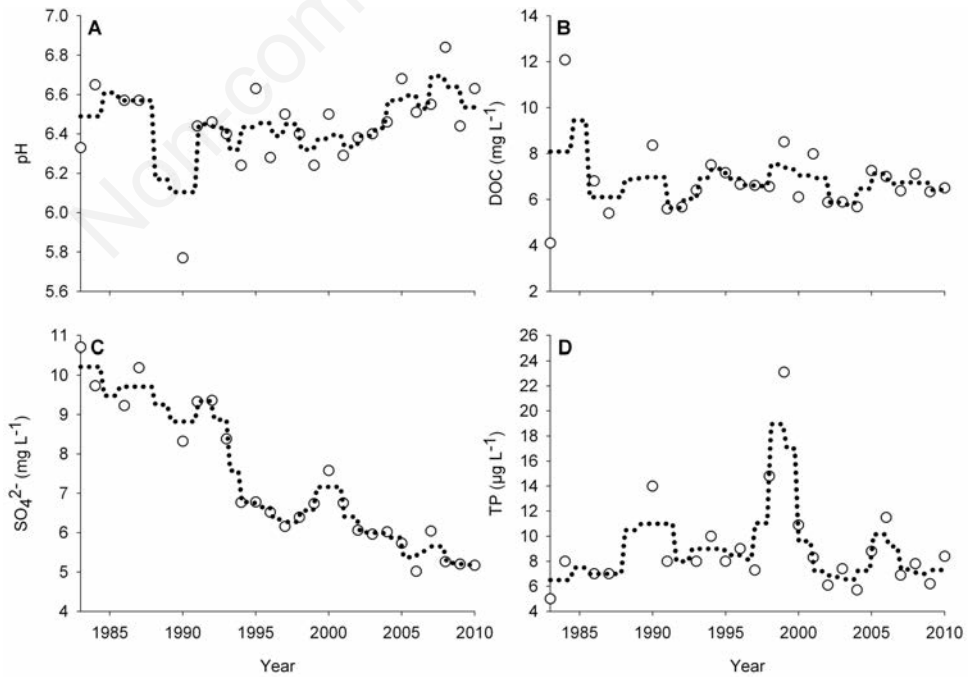


Fig. 2. Selected monitoring data (A, pH; B, DOC; C, SO_4^{2-} ; D, TP) from W16 between 1983 and 2010. Dotted lines represent running average with sampling proportions of 10%. Data are from the Canadian Wildlife Service Acid Rain Biomonitoring Program (Ontario region).

Laboratory methods

Chaoborus mandibles were isolated from W16 sediments following standard methods outlined in Walker (2001). Wet sediments were preferentially used, with weights ranging from 0.78 to 2.80 g. When wet sediments did not yield sufficient numbers of mandibles, additional freeze-dried sediments were analysed. All sediments were deflocculated with 5% KOH solution on a 200°C hot plate for ~20 minutes. Sediments were then poured onto a 100- μm mesh sieve and rinsed with deionised water before being transferred into a beaker. Aliquots of each sample were poured into a Bogorov tray and viewed under a dissecting microscope at ~25X magnification. Each sample was scanned four times to ensure all mandibles were recovered. Forceps were used to remove mandibles and to place them onto a coverslip. Coverslips with mandibles were then permanently mounted onto a slide using Entellan® (Electron Microscopy Sciences, Hatfield, PA, USA). A minimum count of 10 mandibles per interval was targeted to ensure an accurate representation of *Chaoborus* assemblage composition (Quinlan and Smol, 2010). Out of a total of 26 samples, this criterion was achieved 17 times (mandible count summary: min=5, mean=15.7, max=47).

Bosmina (primarily *Bosmina longirostris*-type) remains were prepared following methods outlined in Korhola and Rautio (2001). Briefly, wet sediments were weighed and deflocculated with 5% KOH solution on a 200°C hotplate for ~20 minutes. Sediments were then poured onto a 38- μm mesh sieve, and rinsed with deionised water, then concentrated into a 50-mL vial. Next, three drops of safranin solution and ethanol were added to increase the visibility of remains and as a preservative, respectively. Production of slides involved pipetting 50- μL aliquots of the prepared solution onto slides. The number of aliquots used per slide ranged from three to five, depending on the concentration of *Bosmina* remains. Slides were then mounted using glycerin jelly.

Bosmina size measures were obtained from 26 intervals of the core, generally matching those intervals sampled for *Chaoborus*.

Measurement and taxonomy

Chaoborus mandibles were identified using a Leica DMR light microscope under bright-field illumination at 200X magnification. Identifications were based upon a taxonomic key by Uutala (1990). The concentration of chaoborids was calculated by enumerating the number of mandibles found for each species per 1 g of dry sediment weight. Student's *t*-test was used to compare average chaoborid concentration before and after the mid-1980s minnow introduction.

Measurements of *Bosmina* antennules, carapaces, and mucros were carried out using a Leica DMR microscope under bright-field illumination at 200X magnification. Slides were scanned completely in a series of horizontal transects and each non-fragmented *Bosmina* carapace, mucro, and antennule was measured according to Korosi *et al.* (2010). Measurements were made using a Retiga digital camera and Northern Eclipse Image Analysis software version 6 (Empix Imaging Inc., Mississauga, Canada). For each interval, at least 40 headshields (with antennules) and 40 carapaces were measured. This is in accordance with Brahney *et al.* (2010), who demonstrated that a minimum of 35 remains measured was necessary for the detection of meaningful variation in *Bosmina* size structure.

Dating

Sediment core samples were analysed for ^{210}Pb and ^{137}Cs activity using gamma spectrometry techniques to establish a chronology for the past ~150 years. Preparation of samples for dating followed methods outlined by Schelske *et al.* (1994). Decays of ^{210}Pb , ^{137}Cs , and ^{214}Bi were measured from 15 sediment intervals. Ages based on unsupported ^{210}Pb concentrations were then calculated

Tab. 1. Numbers of deployed minnow traps and fish collected across years from W16. Data are from the Canadian Wildlife Service Acid Rain Biomonitoring Program (Ontario region).

Year	Number of traps	Northern redbelly dace (<i>Phoxinus eos</i>)	White sucker (<i>Catostomus commersoni</i>)
1983	6	0	0
1986	6	5	0
1989	6	0	0
1990	6	59	2
1994	6	1559	0
1999	2	349	0
2001	6	654	0
2004	6	1591	0
2008	6	1197	0

by the constant rate of supply (CRS) model (Appleby, 2001). The ^{137}Cs peak (a proxy for the height of nuclear fallout prior to the 1963 moratorium on nuclear weapons testing) was then compared to ^{210}Pb -inferred ages as an assessment of our radioisotopic dating efforts (Fig. 3).

Visible reflectance spectroscopy Chlorophyll *a*

To assess temporal patterns in aquatic primary production from W16, sedimentary-inferred Chlorophyll *a* using visible reflectance spectroscopy (VRS Chl-*a*) was measured following Michelutti *et al.* (2005). This method captures the spectral signatures of sedimentary Chlorophyll *a* and associated breakdown products between 650 and 700 nm (Wolfe *et al.*, 2006). In a study designed to test the efficacy of the VRS method, Michelutti *et al.* (2010) showed that relatively accurate trends across lakes of varying productivity and known eutrophication histories were reconstructed. Sediment intervals from W16 were freeze-dried and sieved through a 125- μm mesh sieve, and analysed for VRS Chl-*a* using a FOSS NIRSystem Model 6500 rapid content analyser. VRS Chl-*a* concentration was then calculated using a linear regression equation

from Michelutti *et al.* (2005). We emphasise that it is the trends in VRS Chl-*a* through time, and not the absolute values, that are most informative (Michelutti *et al.*, 2010).

Top-bottom cladoceran assemblage composition

Cladoceran assemblages were enumerated from the top (0-0.25 cm) and bottom (29-29.5 cm) intervals of the W16 core. The top-bottom palaeolimnological approach was chosen as it allows for comparative examination of assemblage changes occurring between modern and pre-industrial times in the lake. Specifically, using this *snapshot* approach we aimed to recognize coarse shifts in the relative abundances of dominant cladoceran groups (*i.e.* benthic vs pelagic taxa). Cladocera were prepared following standard methods outlined in Korhola and Rautio (2001) and enumerated using standard fossil cladoceran counting guidelines (Kurek *et al.*, 2010b).

RESULTS

Radiometric dates

Dates were calculated from the core surface to 19.25

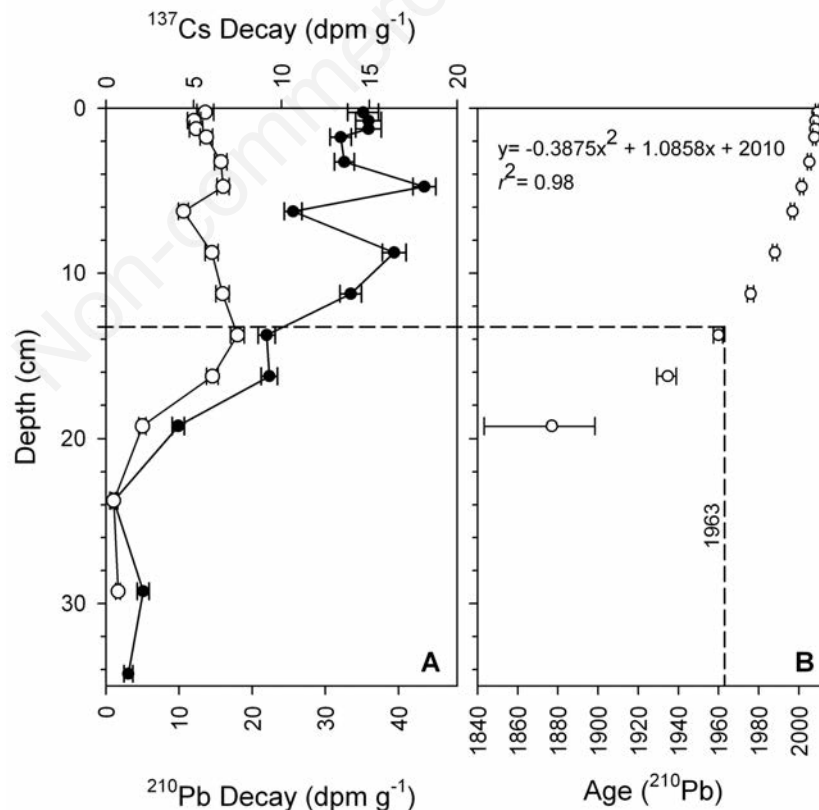


Fig. 3. A, Activity of ^{210}Pb (black circles) and ^{137}Cs (white circles) vs sediment depth for lake W16. B, Plot of ^{210}Pb generated age vs sediment depth for lake W16. Error bars represent standard error. Background levels were reached after a depth of 19.25 cm. At depths greater than 19.25 cm, a second-order polynomial regression was used to estimate ages.

cm, when ^{210}Pb activity reached background levels (Fig. 3A). Uncertainty within the dates increases with depth from approximately ± 1 year in surface sediments to ± 15 years at 19.25 cm. Dates were extrapolated using a second-order polynomial regression (Fig. 3B) for the remainder of the core, with 36.25 cm (core bottom) roughly corresponding to \sim A.D. 1500. The ^{210}Pb -inferred date of 1963 corresponded closely with the ^{137}Cs peak at ~ 13 cm.

Visible reflectance spectroscopy Chlorophyll *a*

Visible reflectance spectroscopy Chl-*a* showed stable and high levels from the base of the core to a depth of ~ 20 cm (corresponding to ~ 1880) (Fig. 4). Visible reflectance spectroscopy Chl-*a* levels from the basal sediments to ~ 20 cm averaged 0.10 mg g^{-1} dry weight (dwt) and were on average the highest in the entire record. Then, they declined abruptly until ~ 17 cm (~ 1920), reaching a low of 0.05 mg g^{-1} dwt. Levels of VRS Chl-*a* then remained constant at an average of 0.05 mg g^{-1} dwt until a depth of ~ 10 cm (corresponding to the approximate time of minnow introduction). After this point, levels showed a steady increase, reaching 0.07 mg g^{-1} dwt in the modern intervals of sediment. These were the highest VRS Chl-*a* values recorded since ~ 1900 .

Chaoborus assemblages

Mandible concentrations were substantially more variable below 10.75 cm (pre- ~ 1980 s) than in the modern sediments from 9.75 cm to present (Fig. 4). A significantly higher average concentration of mandibles occurred in intervals before minnow introduction, compared to those after minnow introduction. ($t=-2.70$, $P=0.01$, $df=24$). Additionally, mandible concentration showed greater variability in lower sediment intervals deposited before minnows were introduced to W16 (Fig. 4). Larger-bodied chaoborids dominated the sediment record and on average composed $\sim 90\%$ of the assemblage (Fig. 4). Concentration of large-bodied chaoborid species decreased and remained at lower levels after the introduction of minnows. Although they appeared several times in the sediment record, smaller-bodied species did not reach the high pre-minnow-introduction concentrations achieved by the larger-bodied species.

Bosmina size structure

Antennule and carapace lengths remained stable throughout the sediment record, whereas average mucro lengths varied significantly (Fig. 4). Mucro lengths were stable and the highest in the entire record from the bottom of the core to approximately 20 cm (~ 1880), averaging 70

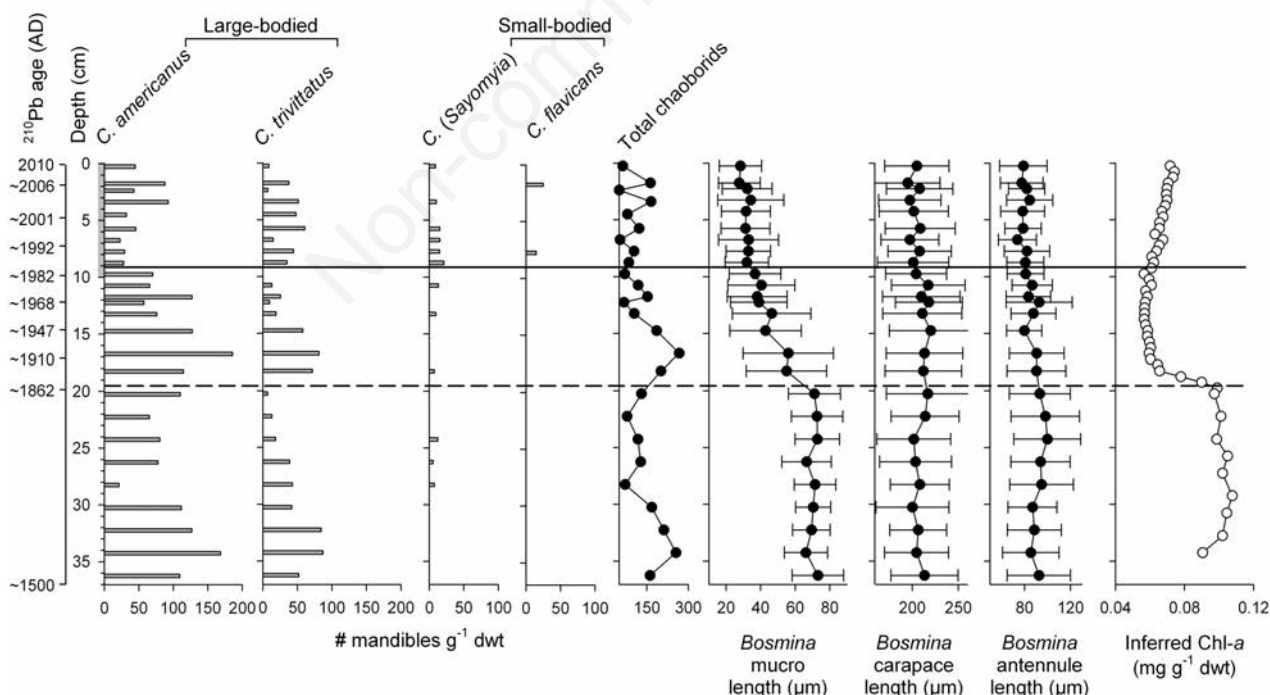


Fig. 4. Stratigraphy of concentrations of individual *Chaoborus* taxa, total chaoborid concentrations, *Bosmina* size measures (error bars show one standard deviation), and sedimentary-inferred VRS Chlorophyll *a*. The solid horizontal line marks the ~ 1980 s minnow introduction. The dashed horizontal line denotes the late-1800s shift in both VRS Chl-*a* and mucro length. The shaded y-axis between 10 and 0 cm represents the period of water chemistry and fish community monitoring at W16.

μm . Mucro lengths declined by $\sim 40 \mu\text{m}$ during the period between ~ 1880 and ~ 1983 . During the period of decline, standard deviation values were comparatively large, indicating a wide range of mucro lengths existed in these intervals. After the minnow introduction, mucro lengths remained stable and values were the lowest in the record, at an average of $31 \mu\text{m}$, for the remainder of the core.

Top-bottom (before and after) palaeolimnological analysis of cladoceran assemblages

Cladoceran assemblages showed notable changes in the relative abundances of pelagic and littoral taxa (Fig. 5). Through the top-bottom palaeolimnological approach, we observed that pelagic taxa increased relative to littoral taxa. A $\sim 36\%$ increase in relative abundance from pre-industrial to modern times was observed in the mainly pelagic taxon *Bosmina*. *Daphnia* spp. were never a major contributor to the assemblage and occurred at only $\sim 1\%$ abundance in both time periods. The dominant littoral taxa in our record, *Alona rustica* and *Chydorus brevilabris*, decreased in modern times by ~ 25 and $\sim 4\%$, respectively.

DISCUSSION

Early ecosystem changes at W16 indicative of a late-1800s shift in stratification regime

Beginning in the late-1800s, a large-magnitude reduction in predation intensity occurs at W16 as noted by an av-

erage *Bosmina* mucro length decrease of $\sim 40 \mu\text{m}$ (Fig. 4). Shifts in the average length of *Bosmina* mucros have been interpreted to reflect variation in predation intensity in many palaeolimnological studies (Alexander and Hotchkiss, 2010; Korosi *et al.*, 2013) and modern surveys (Kerfoot, 1975; Black, 1980; Post *et al.*, 1995). Although antennule and carapace lengths may vary with type and intensity of predation, mucro length is often recognised as the most reliable bosminid response to variations in predation (Korosi *et al.*, 2013). Longer mucro length confers an advantage as it makes *Bosmina* difficult to grasp, handle, and consume by gape-limited invertebrates (Kerfoot, 1975), such as *Chaoborus* larvae. Therefore, we infer that this late-1800s decrease in mucro length of the dominant zooplankton grazer at W16 is a strong indication that invertebrate predation intensity on *Bosmina* was greatly reduced. However, we acknowledge that due to poor fossil representation of many invertebrate predators (*e.g.* copepods, predatory cladocerans, surface-feeding beetles and insects) we are unable to account for the full scope of predation dynamics observed at W16. Additionally, we acknowledge that, due to often fragmented or obscured bosminid remains, genus-level taxonomy was used, and this may have masked size changes arising from a shift in *Bosmina* species.

Chaoborus assemblages at W16 display several distinct shifts in concentration and species composition before minnow introduction in the 1980s (Fig. 4). The high,

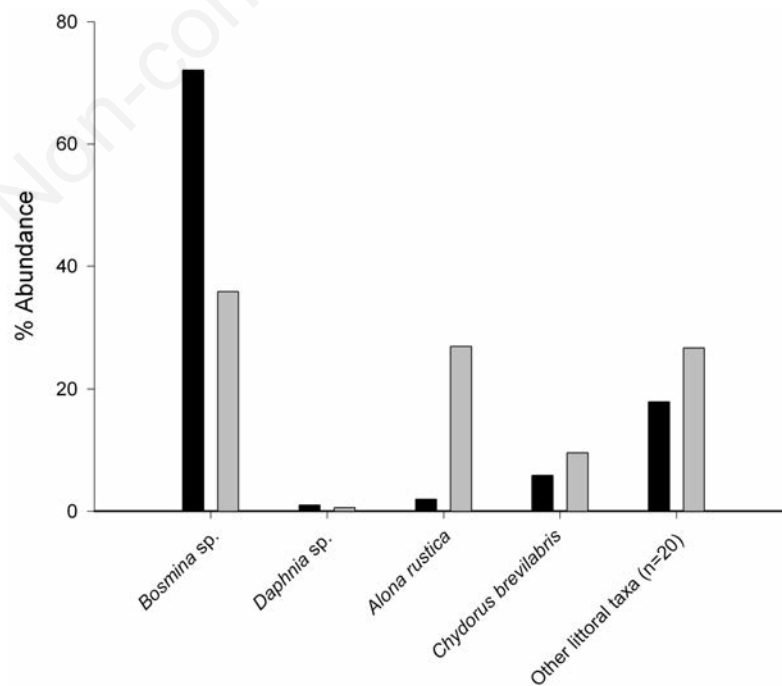


Fig. 5. Relative abundances of cladoceran taxa and groups from top (0-0.5 cm) and bottom (29-29.5 cm) sediment intervals of W16. These intervals represent modern (black bars) and pre-industrial (grey bars) time periods at W16.

but at times variable, concentration of *C. americanus* and *C. trivittatus* mandibles in sediments deposited prior to ~1980 is indicative of very low fish planktivory, as these larger-bodied chaoborids are often heavily preyed upon by planktivorous fish (von Ende, 1979; Sweetman and Smol, 2006). In eastern North America, *C. americanus* has been observed in only a handful of lakes with fish present, and in these instances fish communities showed low richness (Uutala *et al.*, 1994) or *C. americanus* mandibles were sparse (Schilling *et al.*, 2008). Collectively, our *Chaoborus* concentration and assemblage composition data suggest that W16 was historically fishless, or at the very least, only low numbers of non-planktivorous fish were ever present.

The close relationship between reduced invertebrate predation intensity and decreased primary production (inferred by VRS Chl-*a*) (Fig. 4) suggests that a common change in lake functioning at W16 may structure these measures. One mechanism that explains the apparent correlation between these observed trends is a historic shift in the dominant lake stratification regime at W16. Physical properties of lakes (*e.g.* stratification regime, water colour and temperature) play important roles in lake plankton dynamics, nutrient cycling, and community structure, which ultimately influence higher trophic levels and lake functioning (Magnuson *et al.*, 1997; Keller, 2007). In many northwestern Ontario lakes, a longer and warmer ice-free season due to recent air temperature increases, in turn, affects fundamental water-column processes of lakes (Magnuson *et al.*, 1997; Schindler *et al.*, 1996; Keller, 2007; Rühland *et al.*, 2010). Quinlan *et al.* (2012) demonstrate that small lakes with maximum water depths between ~5 and 8 m in northwestern Ontario are highly susceptible to climate-induced shifts in thermal regime, although differences in wind exposure, lake morphometry, and water colour among lakes are also important factors to consider. The small size and ~8 m depth of W16, in addition to its relatively stained water colour, makes this site especially vulnerable to climate-induced changes in its dominant thermal regime (Magnuson *et al.*, 1997; Keller, 2007; Quinlan *et al.*, 2012).

Climate observations and paleoclimate reconstructions from the Great Lakes region of Ontario show that climate during the 20th century, especially the latter decades, largely differs from conditions over the previous centuries (Magnuson *et al.*, 1997; Buhay and Edwards, 1995; Buckley *et al.*, 2004; Rühland *et al.*, 2010). Specifically, dendroclimatological inferences of climatic conditions during the late-1800s were much cooler and drier than those of the 20th century (Buhay and Edwards, 1995; Buckley *et al.*, 2004). Prolonged periods of less precipitation have been shown to reduce DOC export from the catchment, thus influencing water colour and light transparency (Schindler *et al.*, 1996). Evidence of lower lake levels at W16 is suggested by

greater relative abundances of littoral taxa (*i.e.* *A. rustica* and *C. brevilabris*) than pelagic taxa (*i.e.* *Bosmina*) during pre-industrial compared to modern times (Fig. 5). If during the mid-to-late-1800s at W16, cooler air temperatures prevailed, lake depth was lower, and DOC concentrations of lake water were reduced, then W16 may have been less likely to record periods of strong thermal stratification. At the turn of the 20th century, as the regional climate became warmer and wetter (Buhay and Edwards, 1995), a shift in the dominant thermal regime at W16, from a polymictic system that may have only stratified weakly or infrequently, to a lake that experiences longer or more intense periods of thermal stratification, is possible. Therefore, the onset of more regular periods of thermal stratification can explain both the decline in VRS Chl-*a* levels and the reduction in predation intensity, as well as the close synchronicity between these two events.

Using enclosure experiments in a small, shallow fishless lake from Sudbury, MacPhee *et al.* (2011) demonstrated that the dominant lake stratification regime and overall water-column temperatures affect *Chaoborus* predation rates and success. Specifically, it was noted that under isothermal conditions, *Chaoborus* predation effectively lowered zooplankton abundance; however, under thermally-stratified conditions, *Chaoborus* did not exhibit a significant predation effect. In thermally stratified lakes, zooplankton that show DVM behaviour are likely to experience at least some periods of cooler water temperatures, compared to zooplankton within nearby isothermal lakes. Cooler water temperatures decrease rates of digestion, metabolism, and respiration (Swift, 1976; Giguère, 1981; Spitze, 1985), all of which often increase prey-handling time, and consequently reduce predation rate and success (MacPhee *et al.*, 2011). Hence, predation intensity as inferred by *Bosmina* size attributes may be sensitive to climate-induced shifts in lake water properties.

In contrast to the palaeolimnological-based findings of Korosi *et al.* (2012), which demonstrate that *Daphnia* may control primary production through grazing, we note a large decline in VRS Chl-*a* levels, while *Daphnia* abundances remain exceptionally stable and low (Figs. 4 and 5). Our top-bottom analysis shows no change in *Daphnia* relative abundance and a large (~36%) increase in relative abundance of the smaller-bodied *Bosmina* between pre-industrial and modern times (Fig. 5). Based on these observed trends, we conclude that grazing by *Daphnia* were not responsible for the decline in VRS Chl-*a* levels at W16 beginning in the late-1800s. Again, our data indicate that both primary production and predation intensity were influenced by a shift in the dominant thermal regime of W16. Specifically, after the late-1800s, W16 likely experienced greater frequency, duration, and intensity of thermal stratification due to regional climatic warming (Buhay and Edwards, 1995; Magnuson *et al.*, 1997), com-

pared to earlier time periods as a primarily polymictic lake. Within polymictic shallow lakes, internal nutrient distribution is often relatively uniform throughout the water column and therefore can support higher levels of primary production (Wetzel, 2001; Søndergaard *et al.*, 2003). However, once a lake thermally stratifies, phosphorus is largely prevented from being entrained from the hypolimnion to the epilimnion for longer periods of time, often resulting in lower levels of overall primary production (Wetzel, 2001). If W16 experienced a historic transition from regular mixing to more frequent or longer periods of thermal stratification, reduced overall primary production in this dystrophic lake is a possible outcome.

Ecosystem changes at W16 after the 1980s minnow introduction

The average minimum concentration of *Chaoborus* mandibles is reached during the mid-1980s, corresponding with the timing of minnow introduction at W16. The minnow introduction did not cause the initial decline in *Chaoborus* concentration at ~1950, but increased numbers of minnows since the early 1990s (Tab. 1) likely stabilised the chaoborid population at typically the lowest levels observed from the W16 sediment record. Because larger-bodied species such as *C. americanus* and *C. trivittatus* can exclude smaller-bodied chaoborids, minnow predation on the more visible, larger-bodied species may have eased predation pressure on smaller chaoborid species; hence, their increased representation at W16 following the minnow introduction (Fig. 4). The lack of complete extirpation of *C. americanus*, despite increasing

minnow populations (Tab. 1), is likely due to the diet of the northern redbelly dace, as it feeds on primarily plant material and also invertebrates (Cochran *et al.*, 1988). Similarly, Uutala *et al.* (1994) observed *C. americanus* present in both fishless lakes and lakes that supported less than two fish species (albeit at low abundance). Although *C. americanus* is often extirpated by planktivorous fish species, we note its coexistence with large numbers of northern redbelly dace (Tab. 1).

After minnows became well-established in 1994, *Bosmina* mucro sizes remain at a constant short length, indicating overall low invertebrate predation intensity. Percent abundance of *Bosmina* also shows a massive increase of ~36% from pre-industrial to modern time periods (Fig. 5), suggesting that within-lake conditions have become more favourable for pelagic zooplankton. The introduction of minnows in the 1980s likely resulted in some fish predation upon *Chaoborus*, thus releasing *Bosmina* from strong invertebrate predation pressure. Warming annual temperatures in the region (Fig. 6) may also have contributed to an increased rate of growth and reproduction in *Bosmina* (Vijverberg, 1980). Furthermore, a warmer climate enables phytoplankton to grow faster, and for a longer period of time (Rhee and Gotham, 1981), providing an increased food supply for pelagic zooplankton. This is evidenced by increasing levels of VRS Chl-*a* since the mid-1980s, which correlate with rising annual temperatures in the Sudbury region (Fig. 6). The large increase in relative abundance of *Bosmina* was likely caused by a combination of biotic and abiotic factors, which led to its increased success in the lake's more recent history.

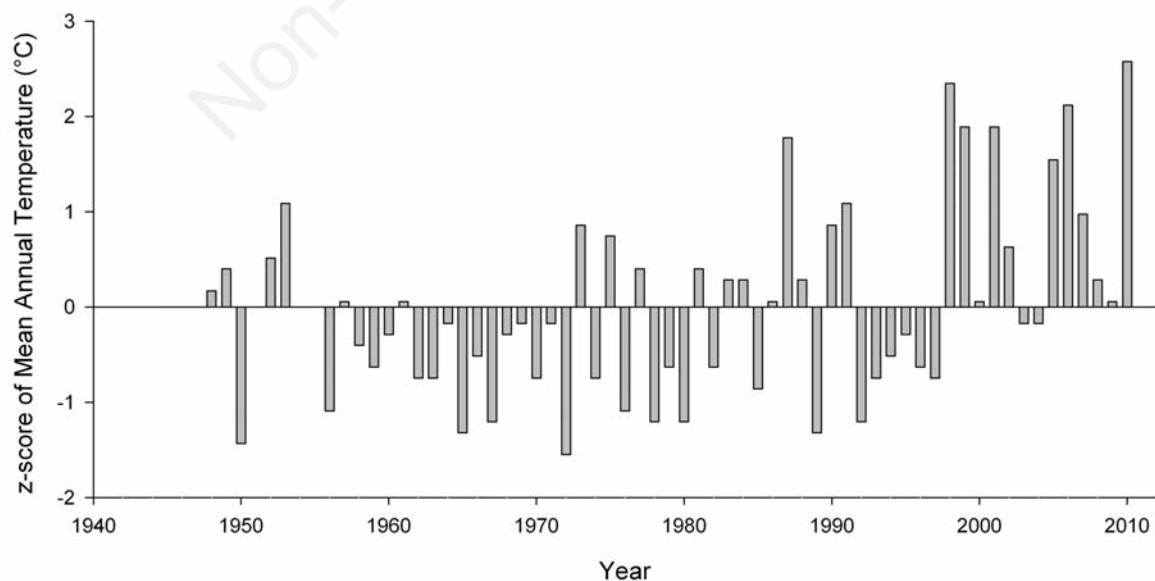


Fig. 6. Standardised values (Z-scores) of annual temperatures for Sudbury, Ontario (climate station 068150). Data obtained from Environment Canada's Adjusted and Homogenized Canadian Climate Records (www.ec.gc.ca/dccha-ahccd/).

Overall, the scale of predation impacts following the 1980s minnow introduction and their subsequent population increase appears negligible compared to the historic shift in predation intensity that began in the late-1800s.

CONCLUSIONS

We noted a large-magnitude decrease in invertebrate predation intensity upon a common zooplankton grazer, *Bosmina*, in the late-1800s at lake W16. Nearly synchronous with the decrease in average *Bosmina* mucro length, VRS-inferred Chl-*a* concentrations indicated an abrupt decline in primary production. Collectively, our multi-proxy data suggest that W16 experienced a historic shift in its dominant thermal stratification regime, subsequently reducing *Chaoborus* predation upon *Bosmina*, and also impacting overall lake primary production. Additionally, we observed that *C. americanus* was not extirpated following the 1980s minnow introduction at W16. Although the presence of *C. americanus* is often interpreted as indicating fishless conditions, its coexistence with a growing population of minnows at W16 provides strong, observational evidence that care should be taken when using indirect measures alone to define fish introduction or extirpation from lake sediment records. Because many small, shallow lakes are fishless and thus support distinct interactions among the biota present, as well as are vulnerable to climate-induced shifts in thermal regime, their value in understanding long-term ecological change continues to be an active research endeavour.

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